


Clay content and soil moisture mapping using on-ground time-domain GPR

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Soil characterization and Ground Penetrating Radar (GPR)

GPR bridges a scale gap between point scale and remote sensing scale, namely, field scale.



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GPR for soil moisture sensing

- On-ground GPR methods (Huisman *et al.* 2003)
 - Common midpoint (CMP)
 - Wide angle reflection and refraction (WARR)
 - Common offset method for Single Trace Analysis (STA)



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GPR for soil moisture sensing


- Off-ground GPR methods (Lambot *et al.* 2007)
 - Surface reflection coefficient
 - Full wave-form inversion using far-field GPR model (Lambot *et al.* 2004)

$$R = \frac{1 - \sqrt{\epsilon_r}}{1 + \sqrt{\epsilon_r}}$$

$$\epsilon_r = \left(\frac{1 - R}{1 + R}\right)^2$$


$$\frac{R}{R_{PEC}} = \frac{\epsilon_r}{\epsilon_{PEC}}$$

$$R_{PEC} = -1$$



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Why on-ground time-domain GPR?

Common-offset method and Direct Ground Wave (DGW)



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The approach description


Hypothesis: The ground is homogeneous around GPR antenna-set

> Dielectric permittivity and soil moisture estimation:

$$\Delta t = \frac{\Delta x}{v}$$

$$v = \frac{c}{\sqrt{\epsilon_r}}$$

$$\epsilon_r = \left(c \cdot \frac{\Delta t}{\Delta x}\right)^2 \text{ Using Topp's model } \theta$$



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Methodology

The approach description

Hypothesis: 2D isotropic radiation

➤ Electrical conductivity estimation:

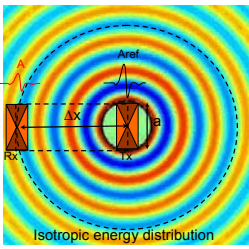
Assuming $\sigma=0$:

$$S_r = \pi \Delta x^2$$

$$S_r = a \Delta x$$

$$\frac{P_1}{P_2} = \frac{A_{ref}^2}{A^2} = \frac{S_r}{S_r} = \frac{\pi \Delta x^2}{a \Delta x} = \frac{\pi}{a} \Delta x$$

$$\Rightarrow \frac{A_{ref}}{A} = \sqrt{\frac{\pi}{a} \Delta x}$$

$$\Rightarrow \frac{A'_1}{A'_2} = \sqrt{\frac{\Delta x_2}{\Delta x_1}}$$


Isotropic energy distribution

Note: The antenna center frequency is relevant to the antenna dimensions therefore this equation can be approximately used for antenna center frequency.

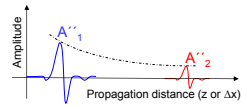
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Methodology

The approach description

➤ Electrical conductivity estimation:

In existence of σ :



$$E_z = E_{z_0} e^{-\alpha z} \cdot \cos(\omega t - \beta z) \hat{a}_z$$

$$\alpha = \omega \sqrt{\frac{\mu \epsilon}{2} \left(1 + \left(\frac{\sigma}{\omega \epsilon} \right)^2 - 1 \right)}$$

Assuming the simple TEM wave propagation for x-directed electric field in homogeneous medium:

$$E \propto \text{Amplitude} \Rightarrow e^{-\alpha(\Delta x_2 - \Delta x_1)} = \frac{A'_2}{A'_1}$$

$$A = A' A' \Rightarrow \frac{A_2}{A_1} = e^{-\alpha(\Delta x_2 - \Delta x_1)} \sqrt{\frac{\Delta x_2}{\Delta x_1}}$$

$$\Rightarrow \alpha = \frac{-\ln\left(\frac{A_2}{A_1} \sqrt{\frac{\Delta x_1}{\Delta x_2}}\right)}{\Delta x_2 - \Delta x_1}$$

$$\Rightarrow \sigma = \omega \epsilon \left(\frac{2 \left(\frac{\Delta x_2}{\Delta x_1} \right)^2}{\left(\frac{A_2}{A_1} \sqrt{\frac{\Delta x_1}{\Delta x_2}} \right)^2 - 1} - 1 \right)$$

How mixing the frequency domain equation and time domain information???

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Methodology

The approach description

Hypothesis: Linear relationship between aggregate electrical conductivity and Clay content

➤ Clay content estimation:

Quadratic parallel model: $\sqrt{\sigma_e} = \theta \sqrt{\sigma_w} + (1-\theta) \sqrt{\sigma_{ag}}$ (Tabbagh et al. 2006)

Linear model: $\sigma_e = \theta \sigma_w + (1-\theta) \sigma_{ag}$

Clay content model: $C_{clay} \% = a \sigma_{ag} + b$

Water conductivity=0.05 S/m

Water content

Apparent conductivity derived by GPR

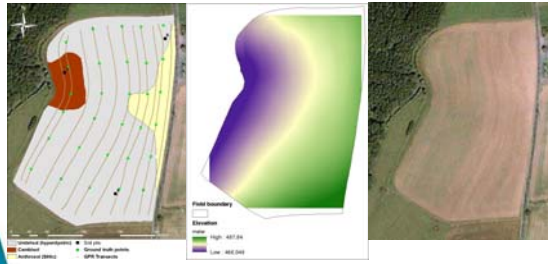
Aggregate conductivity is assumed to be a function of hydraulic conductivity which is dependent of the capillary size

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Study area

The study area

➤ A ~5 ha agricultural field located in the Oëling hills in North part of Grand Duchy of Luxembourg with a mean altitude of 480 m.




Field boundary map
Topographic map derived by dGPS
Google map

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Soil core sampling

The ground-truths and the land conditions

- 30 locations entire the field
- 4 different depths (0-10, 10-20, 20-30, and 30-40 cm) for soil moisture validating.
- 2 different depths (0-10, and 20-30 cm) for clay content validation (using Coulter counter, the cost is ~7€ per sample).

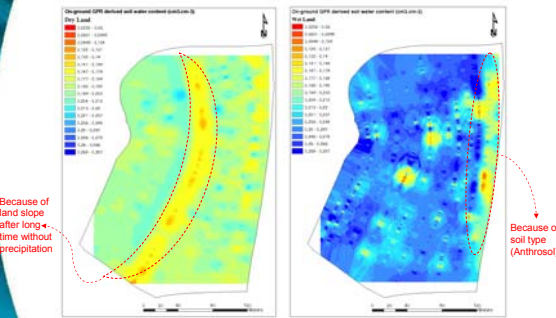


Note: because of unforeseen precipitation event, we had 2 different land conditions namely dry and wet. We applied GPR for both conditions but only 3 locations of ground-truths belonged to the dry land and the rest belonged to the wet land.

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Results

The water content maps



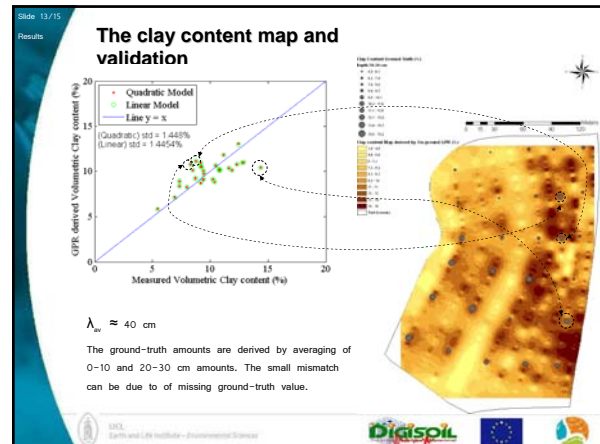
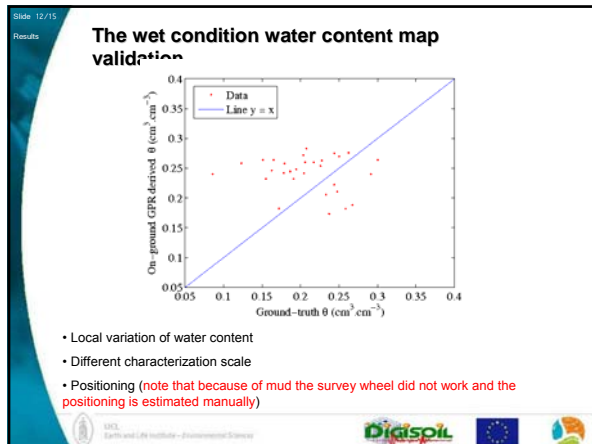
Original GPR derived soil water content (0-10cm)

Re-ground GPR derived soil water content (0-10cm)

Because of land slope after long-time without precipitation

Because of soil type (Anthrosol)

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Conclusions

Conclusions

- Both water content maps derived by on-ground GPR are consistent and show the ability of common-offset method for estimating soil water content.
- The DGW penetration depth is about GPR wave-length which is a function of soil moisture and GPR center frequency.
- There is no significant benefit using linear or quadratic parallel model in order to retrieve the aggregate electrical conductivity from the apparent soil electrical conductivity.
- There is a linear relationship between clay content and aggregate electrical conductivity.
- The clay content map derived by GPR satisfies the ground-truths.

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Perspectives

Perspectives

- Evaluating the proposed approach for other GPR antennas with different center frequencies.
- The effect-investigation of different antenna offsets on retrieving the apparent soil electrical conductivity using DGW.
- Modeling the SOC using the aggregate electrical conductivity derived by proposed approach.

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Thanks for your attention

To whom it may interest
Also we are applying the proposed approach to model the SOC. If you are interested on this job, for more details contact me via the following Email address:
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